α -PbO₂-type high-pressure polymorph of GeO₂

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We have studied the high pressure polymorphism of GeO₂ at pressures up to 60 GPa and temperatures to ~ 1800 K in a laser-heated diamond anvil cell. We have synthesized an α -PbO₂-type (space group *Pbcn*) phase of GeO₂ and demonstrated that it is the stable post-CaCl₂-type (space group *Pnnm*) polymorph at pressures above 44 GPa. The α -PbO₂-structured GeO₂, with a bulk modulus of 256(5) GPa, is denser than CaCl₂ type by 1.6% at 60 GPa. Our study shows that group-IV element dioxides (SiO₂, GeO₂, SnO₂, and PbO₂) have a common sequence of high-pressure structural transformations: rutile-type⇒ α -PbO₂-type.

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The great interest in the polymorphism of GeO₂ comes from both the fundamental point of a possible common sequence of high pressure structural transformations of group-IV element dioxides (SiO₂, GeO₂, SnO₂, and PbO₂), and practical application as a material with unique optical properties. There are two well established polymorphs of GeO₂ at ambient pressure: α -quartz and rutile-type structures with fourfold- and sixfold-coordinated germanium ions, respectively. The rutile structure of GeO₂ is stable under ambient conditions and transforms to α -quartz at 1280 K.¹ The high pressure behavior of both the quartz and rutile phases of GeO₂ has been studied intensively for the last three decades because its structure is believed to display the high pressure properties of geologically important silica at relatively moderate pressures.² However, there are inconsistencies in the results of high pressure phase transformations of GeO₂ reported by different groups.

A phase transition of α -quartz-type GeO₂ above 6 GPa has been observed with various techniques: x-ray diffraction, extended x-ray absorption fine structure (EXAFS), and infra-red and Raman spectroscopy.^{3–13} It has been established that transformation results in an increase of the Ge coordination number with pressure from 4 to 6. The amorphous structure of this high pressure phase of α -quartz-type GeO₂ above 6 GPa has been reported by several groups.^{3–7} Theoretical molecular dynamic studies indicate that under hydrostatic compression a pressure-induced amorphization of quartz-type GeO₂ originates from the mechanical instability of the quartz lattice.^{8,9} Meanwhile, a crystalline high pressure phase of α -quartz-type GeO₂ with an undetermined structure was observed by several investigators^{10,11} and later a monoclinic structure (space group $P2_1/c$, Z=6) of this phase was reported.^{12,13} The post α -quartz GeO₂ monoclinic-type phase was found to be metastable up to 50 GPa at room temperature.13

Recent *ab initio* calculations predict following sequence of phase transformations of GeO₂: rutile \Rightarrow (19 GPa) CaCl₂ \Rightarrow (36 GPa) α -PbO₂ \Rightarrow (65.5 GPa)*Pa*3 (pyrite).¹⁴ Such phase transitions were either observed, or theoretically predicted for silica SiO₂.^{15–17} The second order transition from rutile- to CaCl₂-type phases of GeO₂ has been reported at 25-27 GPa and room temperature.¹⁸⁻²⁰ An orthorhombic high pressure phase of GeO₂ was synthesized at pressure \sim 50 GPa and 1300 K by Ming and Manghnani.²¹ However, according to Ming and Manghnani²¹ the x-ray diffraction patterns of quenched samples are not the α -PbO₂ phase of GeO₂. Evidence of coexisting rutile and disordered Fe₂N-type structures of GeO₂ up to 32 GPa after heating of vitreous GeO₂ at \sim 1300 K have also been reported.^{21,22} Compression of rutile-structured GeO₂ to \sim 36 GPa and subsequent laser annealing at ~1300 K, as described by Haines et al.,19 also does not result in the transformation of CaCl₂-structured GeO₂ to the α -PbO₂-type phase. Laser heating above 43 GPa of both monoclinic structured (α quartz-type GeO₂ starting material) or CaCl₂-type structured phases (rutile or vitreous glass starting materials) results a mixture of the CaCl₂ and Fe₂N-type phases.¹³ At pressures up to 45 GPa and temperatures up to 2300 K, Ono et al.²⁰ observed only the CaCl₂-type post-rutile phase. In summary, despite extensive studies of GeO₂, a high pressure polymorphism of GeO₂ remains unclear. No evidence of the theoretically predicted α -PbO₂ phase has been reported. In this paper we report in situ x-ray studies of phase transformations of GeO₂ at high pressures and temperatures, direct synthesis of α -PbO₂-type GeO₂ from the CaCl₂-structured phase, and structural refinement of GeO₂ polymorphs up to 60 GPa.

In our high pressure experiments α - quartz-type GeO₂ powder (Alfa, purity 99.9999%) was used. The cell constants (space group $P3_22l$, Z=3) were found to be a = 4.9862(8) Å and c = 5.6440(4) Å, in good agreement with literature data. Various types of diamond anvil cells (membrane, symmetrical, and four-pin) were used for pressure generation up to ~60 GPa. The samples were loaded in the 100–150- μ m holes in stainless steel or rhenium gaskets preindented to a thickness of 30–40 μ m. Platinum-black powder (1–3 wt. %) was mixed with the GeO₂ in order to absorb laser radiation for heating and to provide an internal pressure calibrant for high temperature experiments. The

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FIG. 1. (Color online) Integrated x-ray diffraction spectra of various GeO₂ phases collected at different pressure-temperature conditions: (a) starting material α -quartz-type at ambient conditions after loading in a DAC. (b) Monoclinic- $P2_1/c$ -type structure at 34 GPa at room temperature. (c) CaCl₂-type structure at 36.4 GPa after laser heating at 1600±100 K. (d) Mixture of α -PbO₂-type (shown by indices) and CaCl₂-type structures at 44.6 GPa after laser heating at 1600±100 K. (e) Sample recovered from 60 GPa after laser heating at 1800±100 K.

equations of state of Pt or Au, and/or ruby luminescence, were used for pressure measurements in ambient temperature experiments.²³

High resolution angle dispersive x-ray diffraction experiments were performed at the Advanced Photon Source (APS) at Argonne National Laboratory (GSECARS, sector 13, Chicago, III) and the European Synchrotron Radiation Facility (ESRF, ID30 and SNBL beamlines, Grenoble, France) using an on-line image plate or a CCD detector. Details of the experiments are described in our earlier publications.²⁴ At the APS the double-sided laser-heating technique, with a laser spot of 20–40 μ m, was used.²⁵ The temperature was measured with a Kaiser holographic spectrometer coupled with a CCD. The laser heating at ESRF was carried out using a Nd-YAG (yttrium aluminum garnet) laser ($\lambda = 1064$ nm, 17 W).

Examples of diffraction spectra integrated with FIT2D software²⁶ are shown in Fig. 1. The high pressure induced phase transformation of an α -quartz-type structure of GeO₂ to a poorly crystalline monoclinic structure (space group $P2_1/c)^{13}$ phase is detected above 7 GPa [Fig. 1(b)]. This



FIG. 2. (Color online) Ratio of relative intensities of the strongest reflections from α -PbO₂- and CaCl₂-type structures of GeO₂ vs pressure at room temperature (\bullet) and after laser heating (\blacksquare , \blacktriangle) at 1600±100 K (at 60 GPa the temperature was 1800±100 K).

crystalline phase was observed to at least 52 GPa. During laser heating, as the temperature was increased to 1600 ± 100 K, the gradual transformation of the monoclinic phase directly to the orthorhombic structure was observed at 36.4 GPa [Fig. 1(c)]. The high pressure phase has the CaCl₂-type structure (space group *Pnnm*, Z=2). The lattice parameters 36.4 GPa, for example, a=4.2617(4) Å, b at =4.1268(2) Å, and c=2.7817(4) Å, and unit cell volume of 14.73(1) cm³/mol is just slightly smaller than that reported for the same high pressure phase $(14.90(9) \text{ cm}^3/\text{mol} \text{ at } 36)$ GPa) (Ref. 19) obtained by pressurizing rutile-type GeO₂ at quasihydrostatic (methanol-ethanol-water pressure medium) conditions. Detailed results of high-P and -T behaviors of monoclinic and orthorhombic phases will be published elsewhere.

On further compression to 41 GPa and laser heating at ~ 1600 K the reflections from a new high pressure phase started to grow. The relative intensities of the reflections corresponding to the CaCl₂-structured phase decreased with increasing pressure (Fig. 2). Above 44 GPa [Fig. 1(d)] the diffraction intensities from the new phase are greater than those from the CaCl₂-type phase (Fig. 2), although the coexistence of both phases was observed up to 60 GPa.

The structure of the new phase is an orthorhombic α -PbO₂ type (Fig. 3). The crystal structure of α -PbO₂ (*Pbcn* space group symmetry) can be described in terms of a hcp packing of oxygen with one-half of the available octahedral interstices occupied by germanium ions to form 2×2 zigzag chains of GeO₆ edge-sharing octahedra. In the case of the $CaCl_2$ (*Pnnm*) or rutile ($P4_2/mnm$) structures, germanium ions are arranged in such a way as to generate straight chains of edge-sharing octahedra, which are corner linked to form a three-dimensional network with hcp or distorted hcp arrays of oxygen, respectively.²⁷ The molar volume of α -PbO₂ type is smaller than CaCl₂ type: at 44.5 GPa the difference is 1.4% [14.36(1) and 14.56(5) cm³/mol, respectively], and at 60 GPa it is 1.6% [13.89(0) and 14.12(1) cm³/mol, respectively]. The molar volume versus pressure data were fitted using a third-order Birch-Murnaghan equation of state with



FIG. 3. (Color online) Rietveld full profile refinement of the α -PbO₂-type phase of GeO₂ at 60 GPa after annealing at 1800 \pm 100 K.

K' fixed to 4.²⁸ The fit gave values of the bulk modulus for α -PbO₂ and rutile-type phases of 256(5) and 247(3) GPa, respectively, in good agreement with bulk modulus of 250(9) GPa for the rutile-type phase of GeO₂ reported by Haines.¹⁹

To determine the relative phase stability of the α -PbO₂-type structure, a reversal experiment was conducted, and the pressure was released from 44.5 to 41.3 GPa. After laser heating at ~1600 K the relative intensity of α -PbO₂ reflections decreased by a factor of ~15 (Fig. 2). In a separate experiment, heating of the monoclinic $P2_1/c$ -type phase of GeO₂ at 39.5 GPa and \sim 1600 K results in almost the same ratio of α -PbO₂ to CaCl₂-type phases as for the decompressed and heated one at 41.3 GPa. Compression at room temperature of the sample heated at 41 GPa results in the growth of the α -PbO₂-type phase (Fig. 2). Laser heating at 60 GPa a initiated significant increase in the amount of the α -PbO₂-structured GeO₂ (Fig. 2). On decompression at room temperature from 60 to 29 GPa the proportion of CaCl₂- and α -PbO₂-type GeO₂ was not changed, but transformation of a CaCl₂- to a rutile-type structure was observed at ~ 27 GPa. At ambient pressure, a mixture of rutile-type and α -PbO₂-type structures was detected in almost equal proportions [Fig. 1(e)] with molar volumes of 16.65(5) and 16.41(8) cm³/mol, respectively.

Though a theoretical calculation¹⁴ predicts a single stable phase in a pressure range of 36–65 GPa, we have observed the coexistence of both CaCl₂- and α -PbO₂-type structured GeO₂ up to 60 GPa [Fig. 1(d)]. The nonhydrostatic conditions and temperature gradients in laser-heated diamond anvil cells (DACs) could be the one of reasons for incomplete transformation to α -PbO₂-type GeO₂, especially near the diamond surface, where the temperature can drop by a factor of 5.^{29,30} In contrast to the rutile-CaCl₂-type second order phase transition, when tetragonal symmetry breaks due to distortions of O-Ge-O bond angles, but the Ge ions reside at their previous equilibrium positions, the CaCl₂ $\Rightarrow \alpha$ -PbO₂ phase transformation is accompanied by a large displacement of the cations. The alignment of the octahedra along the **c**



FIG. 4. (Color online) Phase pressure boundary vs metal ionic radius for the group-IV metal dioxides (the solid lines are guides to the eye). Data for pyrite-type structure of SiO_2 (Ref. 17) and GeO_2 (Ref. 14) are theoretically predicted. Experimental data are from Refs. 31 and 33–38, and this study.

direction is lost, and a c lattice vector is almost doubled with respect to the CaCl₂ structure. The higher density of the α -PbO₂ structure (compared to the *Pnnm*-type arrangement) allows a greater oxygen-oxygen separation.¹⁷ The finite volume difference (1.4% at 44.5 GPa, for example) indicates its first order character. This kind of transformation does not occur readily at ambient temperature, and requires significant activation energy to induce cations displacements or as has been proposed for SnO_2 , with a rutile to α -PbO₂ transition via a $P112_1/a$ intermediate phase formally identical to baddelevite.^{31,32} Our observation of the increasing proportion of the α -PbO₂ type phase with respect to the CaCl₂ structured phase, even at room temperature with increasing pressure, as well as the higher density of α -PbO₂ modification of GeO_2 , lead us to conclude that the stable phase of GeO_2 in the pressure range of 44–60 GPa is the α -PbO₂-type structure. Thus, theoretical simulations¹⁴ of the high pressure behavior of GeO₂ are in good agreement with our experimental results.

The observation of the α -PbO₂ phase for GeO₂ supports a possible common sequence of high pressure induced transformation of group-IV element dioxides. The α -PbO₂-structured phases have been reported for the analogous metal dioxides: SiO₂ (between 40 and 80 GPa, depending on starting material and conditions of high-P and -Ttreatment), $^{33-35}$ SnO₂ (at ~12 GPa), 31,36 and PbO₂ (at 1.3 GPa).³⁷ In the cases of SiO_2 , GeO_2 , and SnO_2 even with temperature treatment at high pressures^{33,36} a wide range of coexistence of orthorhombic CaCl₂- and α -PbO₂-type structures have been observed. Figure 4 illustrates the systematic high pressure behavior of group-IV dioxides that all adopt the rutile-type structure with a precise linear relationship between the ionic radii of the cation and the lattice constants a and $c.^{38}$

In summary, x-ray *in situ* studies show that on compression at room temperature α -quartz-type GeO₂ transforms to a poor crystalline phase with monoclinic structure (space

group $P2_1/c$) at 7 GPa, and this exists at least up to 52 GPa. A temperature induced transformation of the monoclinic phase to the CaCl₂-type (space group *Pnnm*) at pressures up to 36.4 GPa was detected. We experimentally observed the first order phase transformation of the CaCl₂-type structure of GeO₂ to the α -PbO₂-type (space group *Pbcn*) starting at 41 GPa. A mixture of the CaCl₂-type and the α -PbO₂-type were found to exist up to 60 GPa, but the relative intensity of

the α -PbO₂-type reflections significantly increased with increasing pressure and temperature. The α -PbO₂-type structure of GeO₂ is quenchable and denser than the rutile-type by 1.4% [densities are 6.37(0) and 6.27(9) g/cm³, respectively] at ambient pressure. Our study shows that group IV element dioxides (SiO₂, GeO₂, SnO₂, and PbO₂) have a common sequence of high-pressure structural transformations: rutile-type⇒CaCl₂-type⇒ α -PbO₂-type.

- ¹V. G. Hill and L. L. Y. Chang, Am. Mineral. **53**, 1744 (1968).
- ²H. Wang and G. Simmons, J. Geophys. Res. 78, 1262 (1973).
- ³S. Kawasaki, O. Ohtaka, and T. Yamanaka, Phys. Chem. Miner. **20**, 531 (1994).
- ⁴S. Kawasaki, J. Mater. Sci. Lett. **15**, 1860 (1996).
- ⁵T. Yamanaka, T. Shibata, S. Kawasaki, and S. Kume, in *High-Pressure Research*, edited by Y. Syono and M. H. Manghnani (Terrapub, Tokyo, 1992), p. 493.
- ⁶G. H. Wolf, S. Wang, C. A. Herbst, and D. J. Durben, in *High-Pressure Research* (Ref. 5).
- ⁷T. Yamanaka, T. Nagai, and T. Tsuchiya, Z. Kristallogr. **212**, 401 (1997).
- ⁸T. Tsuchiya, T. Yamanaka, and M. Matsui, Phys. Chem. Miner. **25**, 94 (1998).
- ⁹T. Tsuchiya, T. Yamanaka, and M. Matsui, Phys. Chem. Miner. **27**, 149 (2000).
- ¹⁰F. Vannereau, J. P. Itie, A. Polian, G. Calas, J. Petiau, A. Fontaine, and H. Tolentino, High Press. Res. 7, 372 (1991).
- ¹¹L. Liu and T. P. Mernagh, High Temp.-High Press. 24, 13 (1992).
- ¹² V. V. Brazhkin, E. V. Tatyanin, A. G. Lyapin, S. V. Popova, O. B. Tsiok, and D. V. Balitskii, Pis'ma Zh. Eksp. Teor. Fiz. **71**, 424 (2000) [JETP Lett. **71**, 293 (2000)].
- ¹³J. Haines, J. M. Leger, and C. Chateau, Phys. Rev. B **61**, 8701 (2000).
- ¹⁴Z. Lodziana, K. Parlinski, and J. Hafner, Phys. Rev. B 63, 134106 (2001).
- ¹⁵K. J. Kingma, R. E. Cohen, R. J. Hemley, and H. K. Mao, Nature (London) **374**, 243 (1995).
- ¹⁶L. S. Dubrovinsky, S. K. Saxena, P. Lasor, R. Ahuja, O. Eriksson, J. M. Wills, and B. Johansson, Nature (London) **388**, 362 (1997).
- ¹⁷D. M. Teter, R. J. Hemley, G. Kresse, and J. Hafner, Phys. Rev. Lett. **80**, 2145 (1998).
- ¹⁸J. Haines, J. M. Leger, C. Chateau, R. Bini, and L. Utivi, Phys. Rev. B **58**, R2909 (1998).
- ¹⁹J. Haines, J. M. Leger, C. Chateau, and A. S. Pereira, Phys. Chem. Miner. **27**, 575 (2000).

- ²⁰S. Ono, K. Hirose, N. Nishiyama, and M. Isshiki, Am. Mineral. 87, 99 (2002).
- ²¹L. C. Ming and M. H. Manghanani, Phys. Earth Planet. Inter. 33, 26 (1983).
- ²²L. G. Liu, W. A. Basset, and J. Sharry, J. Geophys. Res. 83, 2301 (1978).
- ²³M. Yousuf, Semicond. Semimetals **55**, 381 (1998).
- ²⁴ V. B. Prokopenko, L. S. Dubrovinsky, V. Dmitriev, and H.-P. Weber, J. Alloys Compd. **327**, 87 (2001).
- ²⁵G. Shen, M. Rivers, Y. Wang, and S. Sutton, Rev. Sci. Instrum. 72, 1273 (2001).
- ²⁶A. P. Hammersley, S. O. Svensson, M. Hanfland, A. N. Fitch, and D. Hausermann, High Press. Res. **14**, 235 (1996).
- ²⁷L. S. Dubrovinsky, N. A. Dubrovinskaia, V. Prakapenka, F. Seifert, F. Langenhorst, V. Dmitriev, H.-P. Weber, and T. L. Bihan, High Pressure Res. (to be published).
- ²⁸O. L. Anderson, Equations of State of Solids for Geophysics and Ceramic Science (Oxford University Press, Oxford, 1995).
- ²⁹L. S. Dubrovinsky, P. Lazor, S. K. Saxena, P. Haggkvist, H.-P. Weber, and T. LeBihan, Phys. Chem. Miner. **26**, 539 (1999).
- ³⁰M. Manga and R. Jeanloz, J. Geophys. Res. **102**, 2999 (1997).
- ³¹J. Haines and J. M. Leger, Phys. Rev. B 55, 11 144 (1997).
- ³²A. G. Christy, Acta Crystallogr., Sect. B: Struct. Sci. **49**, 987 (1993).
- ³³L. S. Dubrovinsky, N. A. Dubrovinskaya, S. K. Saxena, F. Tutti, S. Rekhi, T. L. Bihan, G. Shen, and J. Hu, Chem. Phys. Lett. 333, 264 (2001).
- ³⁴T. G. Sharp, A. E. Goresy, B. Wopenka, and M. Chen, Science 284, 1511 (1999).
- ³⁵P. Dera, C. T. Prewitt, N. Z. Boctor, and R. J. Hemley, Am. Mineral. 87, 1018 (2002).
- ³⁶S. Ono, E. Ito, T. Katsura, A. Yoneda, M. J. Walter, S. Urakawa, W. Utsumi, and K. Funakoshi, Phys. Chem. Miner. 27, 618 (2000).
- ³⁷W. B. White, F. Dachille, and R. Roy, J. Am. Ceram. Soc. 44, 170 (1961).
- ³⁸A. A. Bolzan, C. Fong, B. J. Kennedy, and C. J. Howard, Acta Crystallogr., Sect. B: Struct. Sci. 53, 373 (1997).